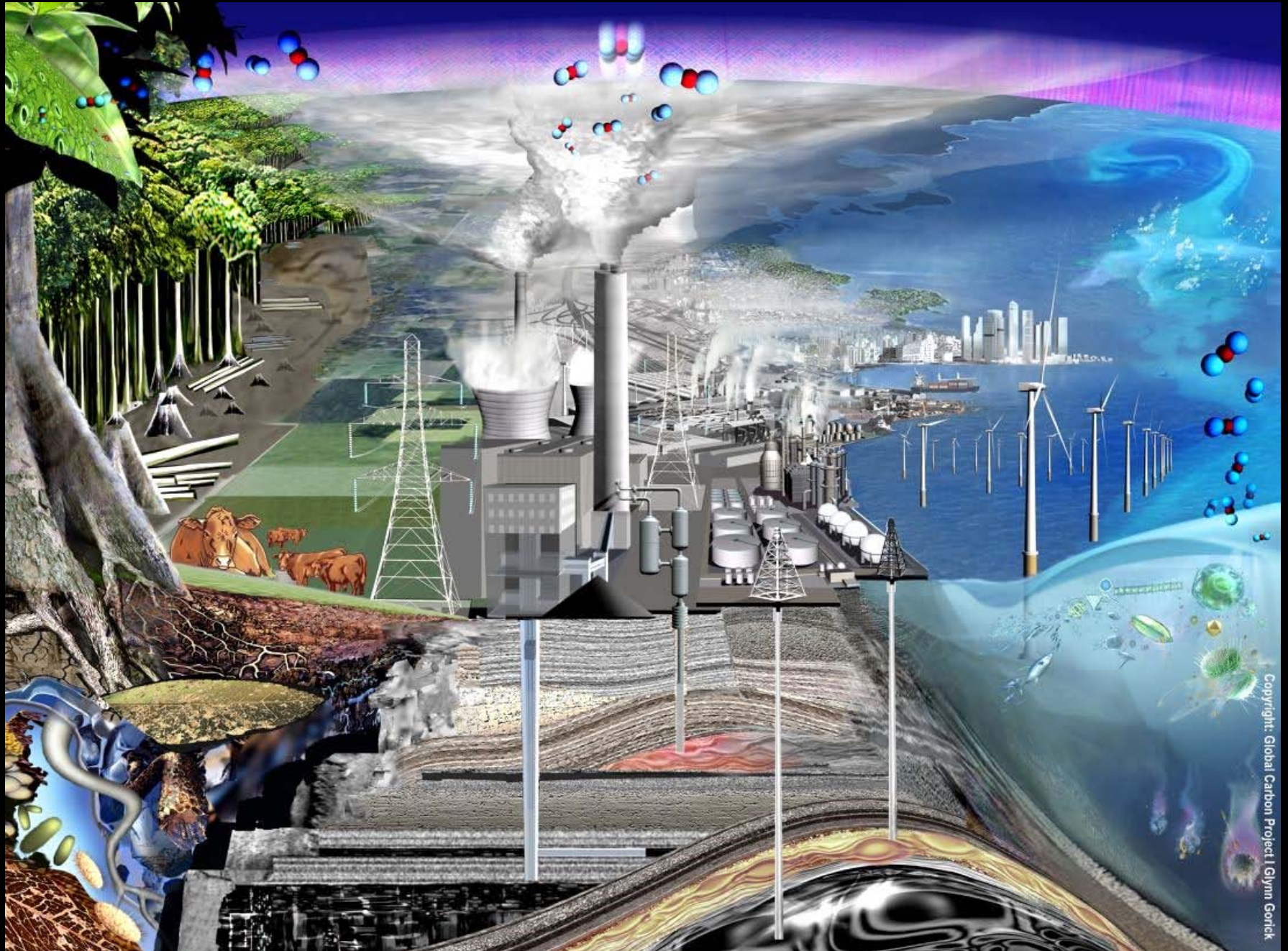


Ocean acidification: a biogeological perspective

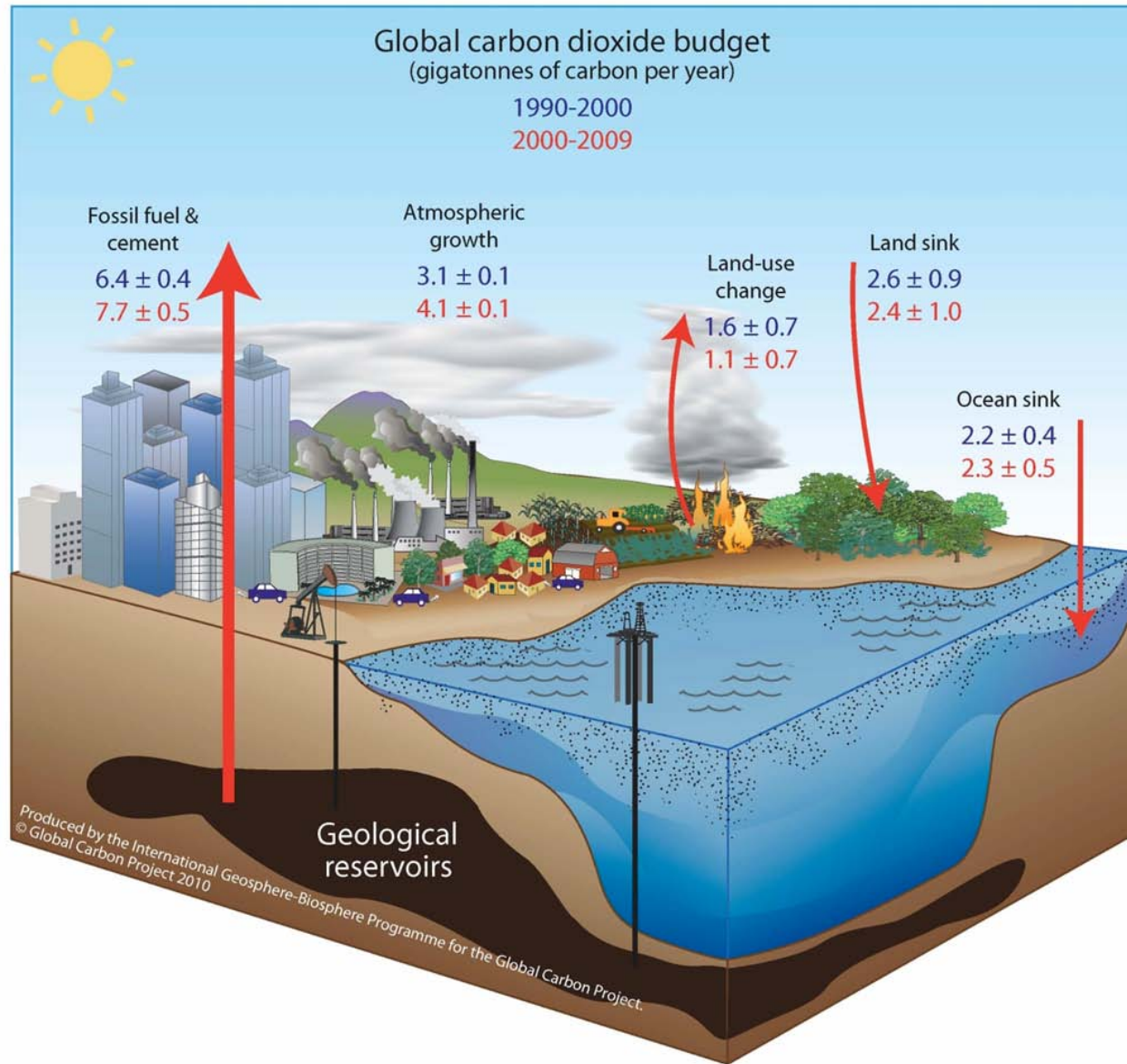
Jelle Bijma (AWI, Bremerhaven, Germany)

- Ocean acidification: present and future
- Why a biogeological perspective?
- Ocean acidification in the past
- Consequences for Biodiversity

Artist Impression of the Human Perturbation of the Carbon Cycle



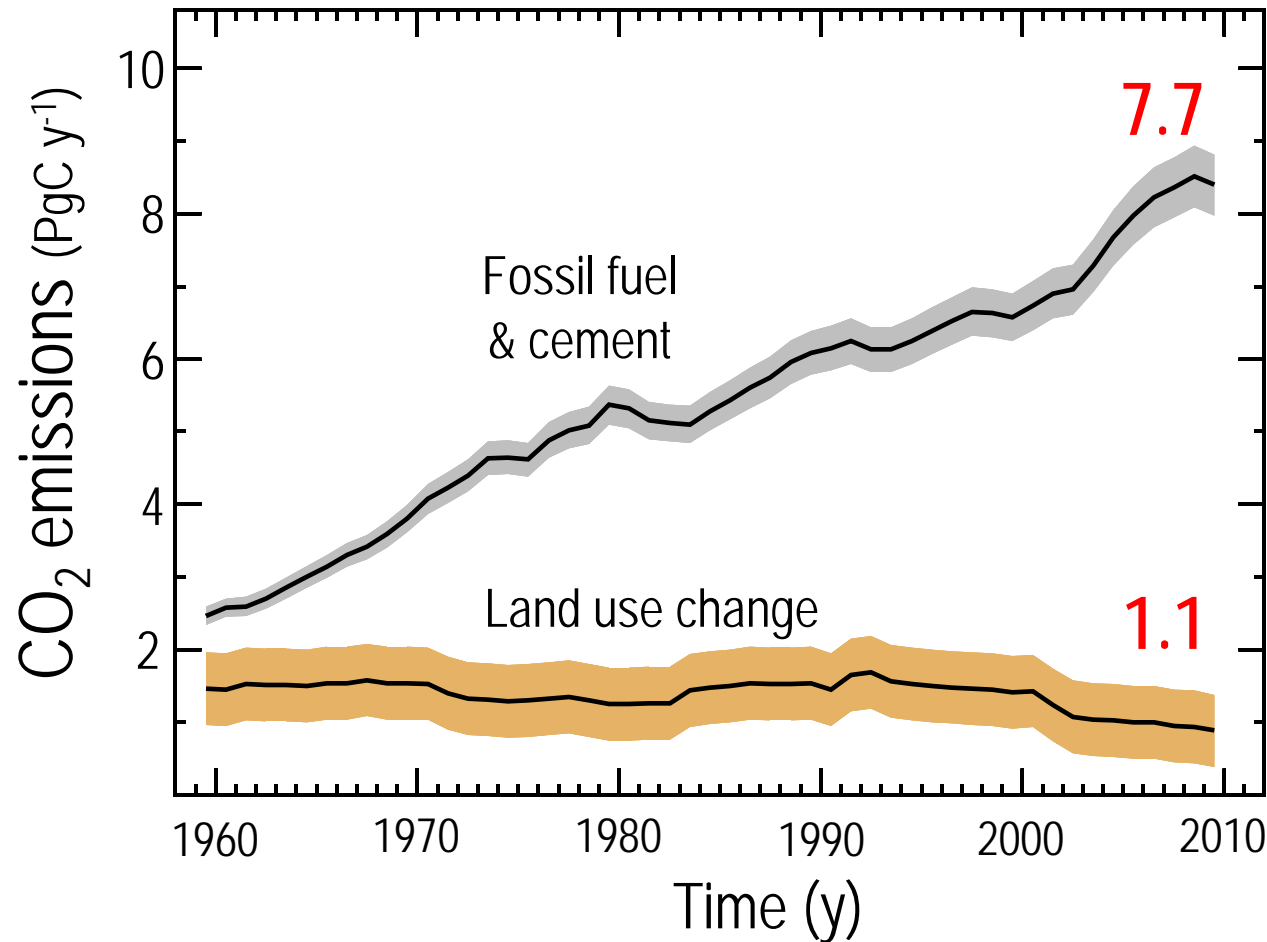
Anthropogenic Global Carbon Dioxide Budget



CO₂ Emissions from Land Use Change (1960-2009)

[1 Pg = 1 Petagram = 1 Billion metric tonnes = 1 Gigatonne = $1 \times 10^{15} \text{g}$]

Average (2000-2009)



8.8 PgC



LUC emissions now
~10% of total CO₂ emissions

Fate of Anthropogenic CO₂ Emissions (2000-2009)

1.1 0.7 PgC y⁻¹



4.1 0.1 PgC y⁻¹

47%



7.7 0.5 PgC y⁻¹ +



4.1 0.1 PgC y⁻¹

47%

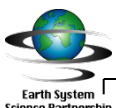
Calculates the residual of all other flux components



26%

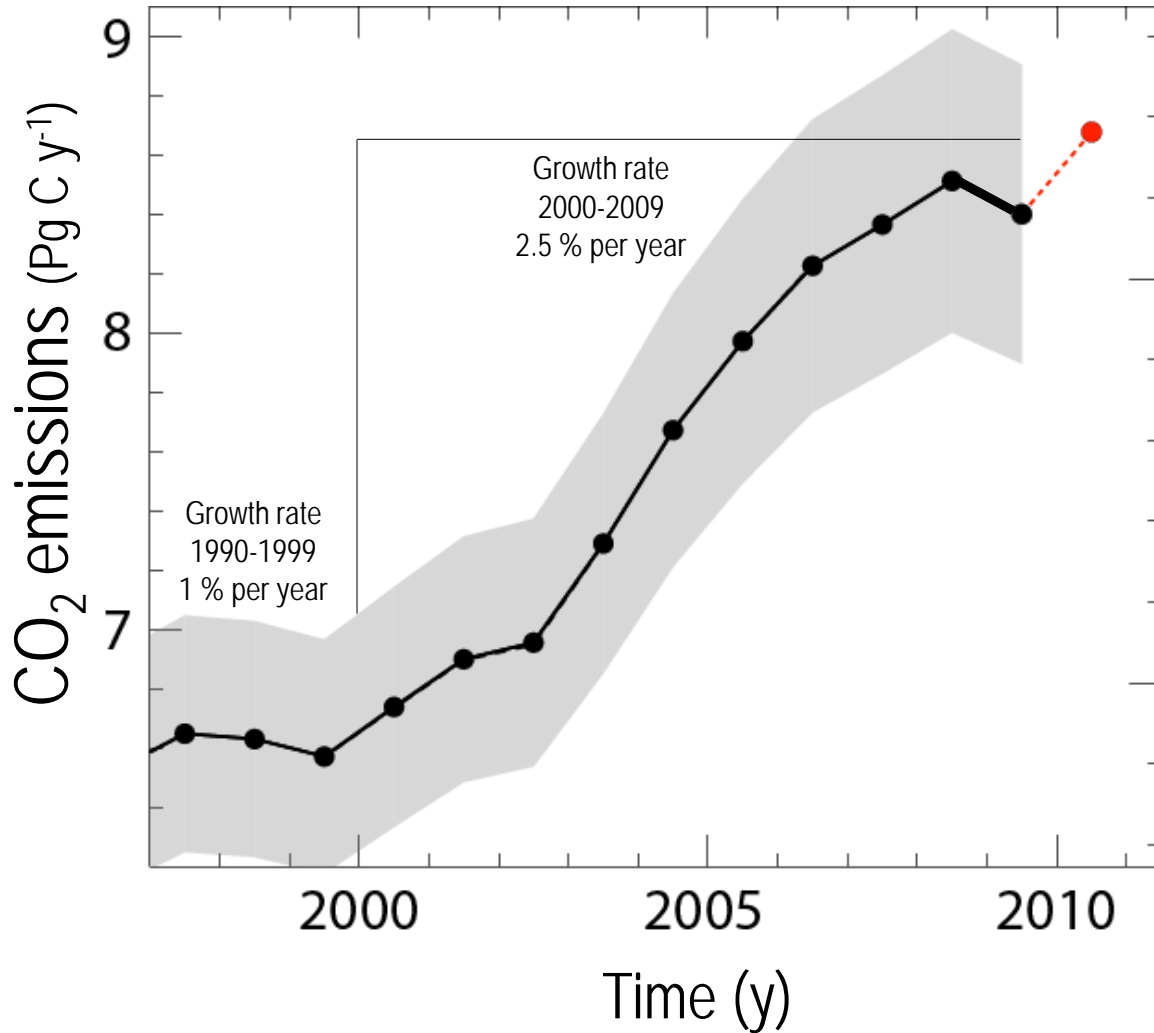
2.3 0.4 PgC y⁻¹

Average of 5 models



Fossil Fuel CO₂ Emissions

[1 Pg = 1 Petagram = 1 Billion metric tonnes = 1 Gigatonne = 1x10¹⁵g]

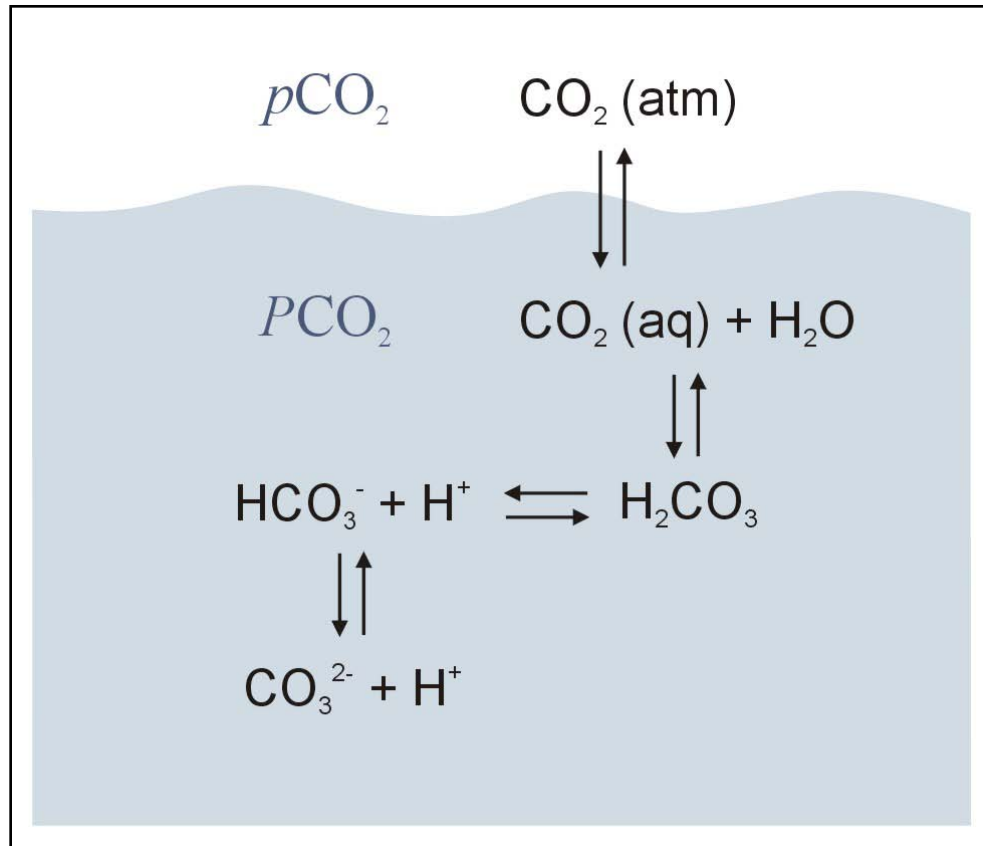


2009:
 Emissions: 8.4 0.5 PgC
 Growth rate: -1.3%
 1990 level: +37%

2000-2008
 Growth rate: +3.2%

2010 (projected):
 Growth rate: >3%

The marine carbonate system



CO_2 (aq): aqueous carbon dioxide

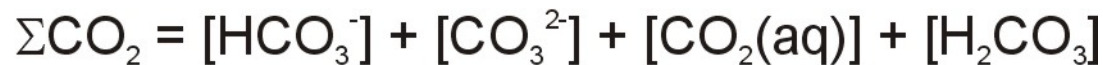
HCO_3^- : bicarbonate ion

CO_3^{2-} : carbonate ion

H_2CO_3 : carbonic acid

ΣCO_2 or DIC or TCO_2 :

Total dissolved inorganic carbon

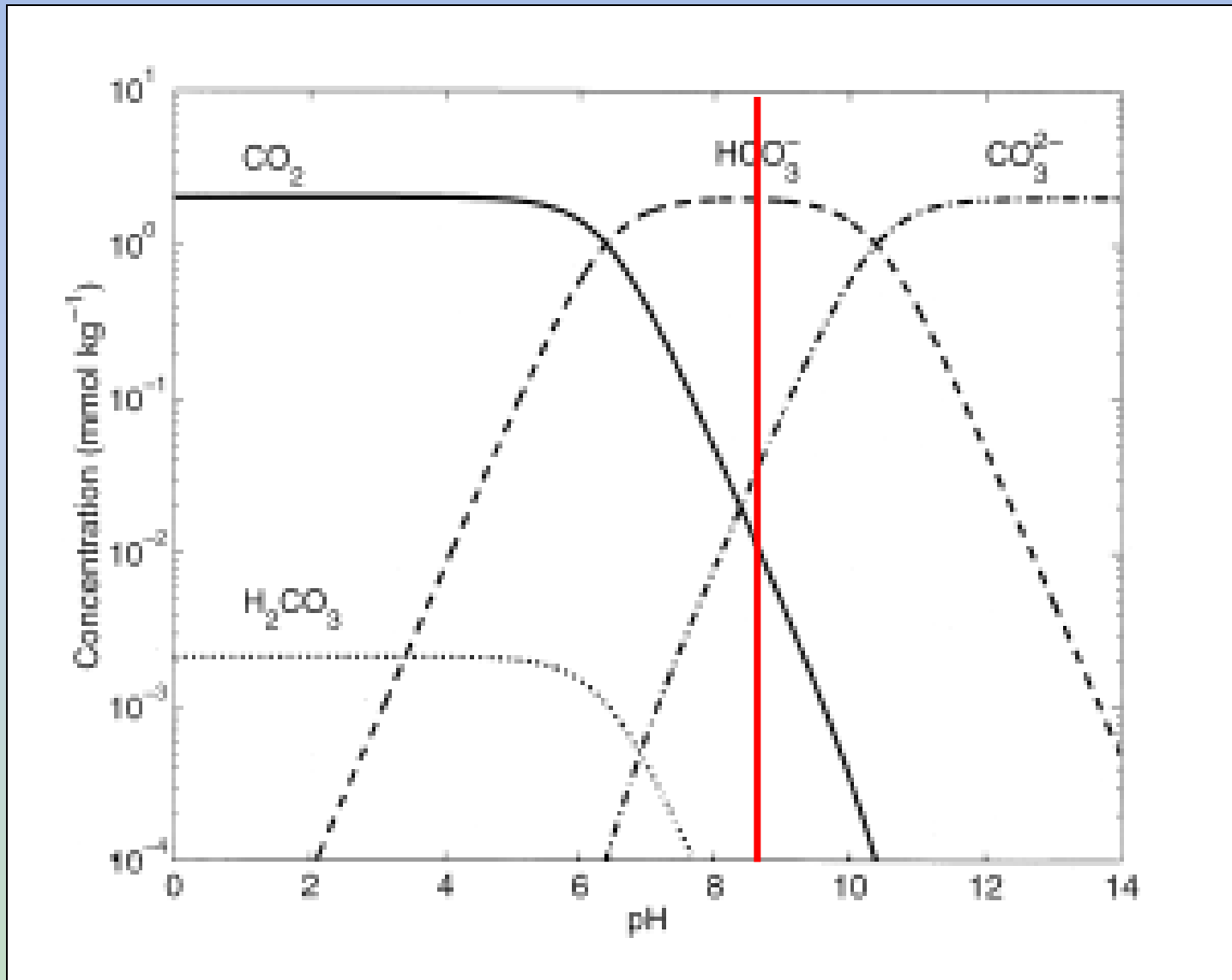


~90%

~10%

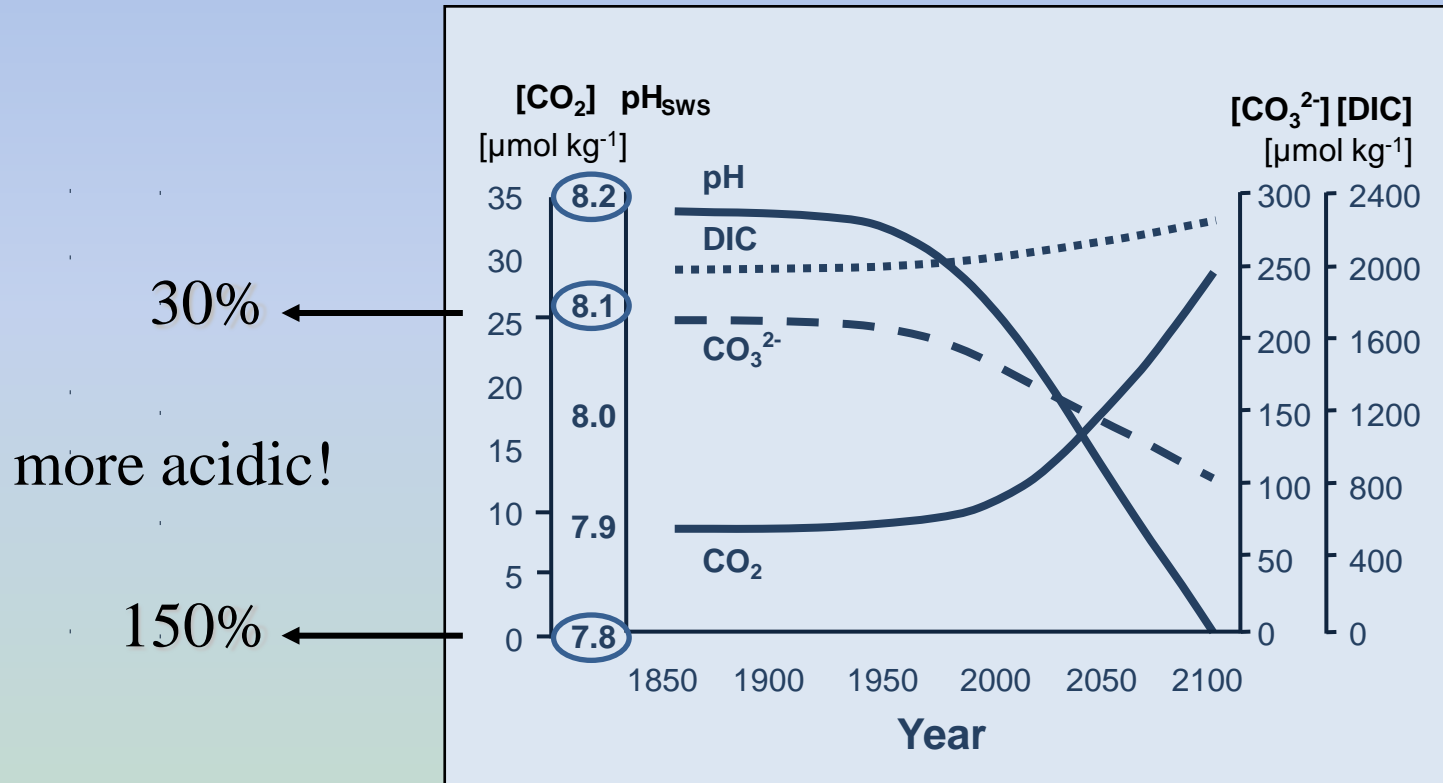
<1%

„Bjerrum plot“

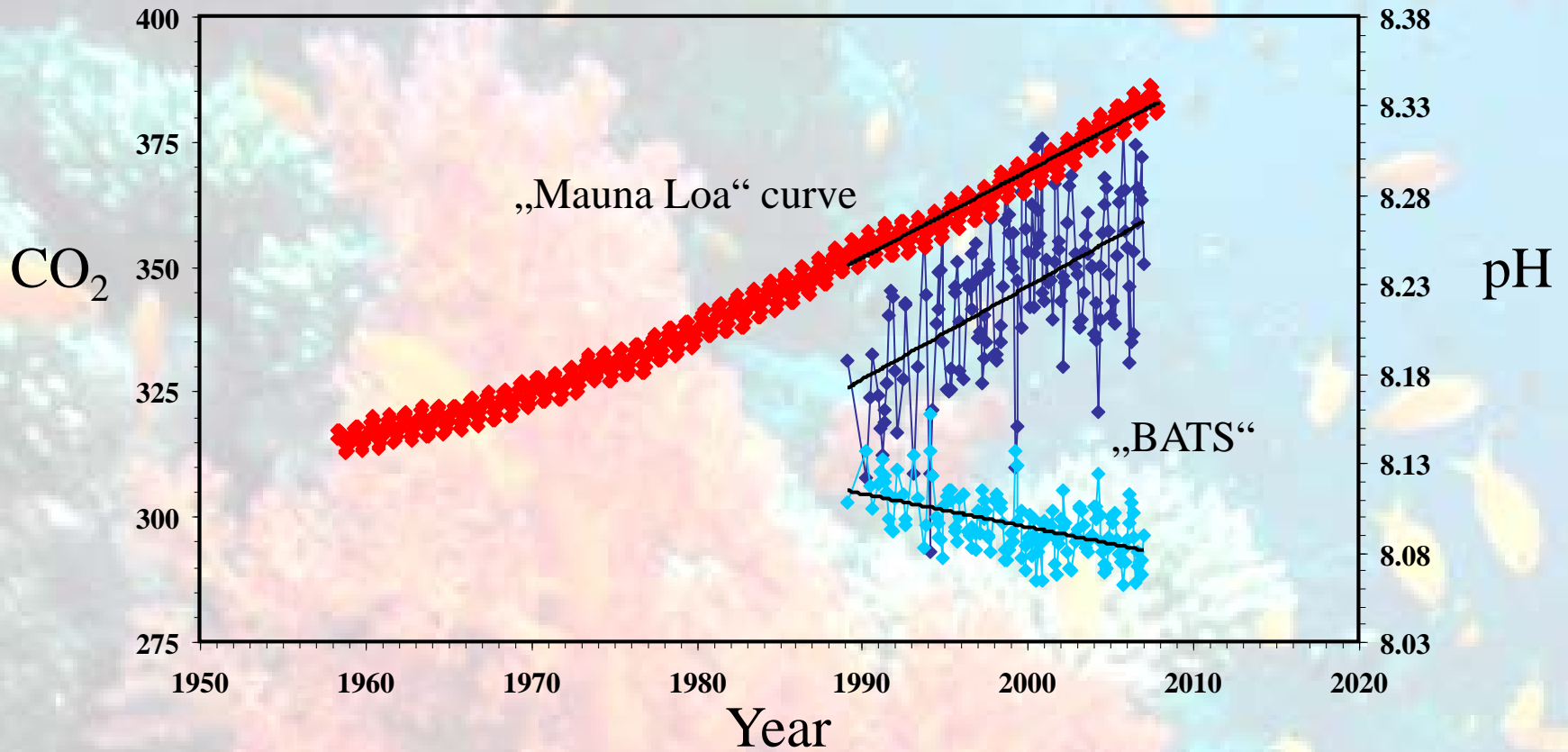


Ocean Acidification

Changes in surface ocean chemistry based on the IS92a scenario IPCC report 1995 (linear increase from 6.3 GtC yr⁻¹ to 20 GtC yr⁻¹ in the year 2100).

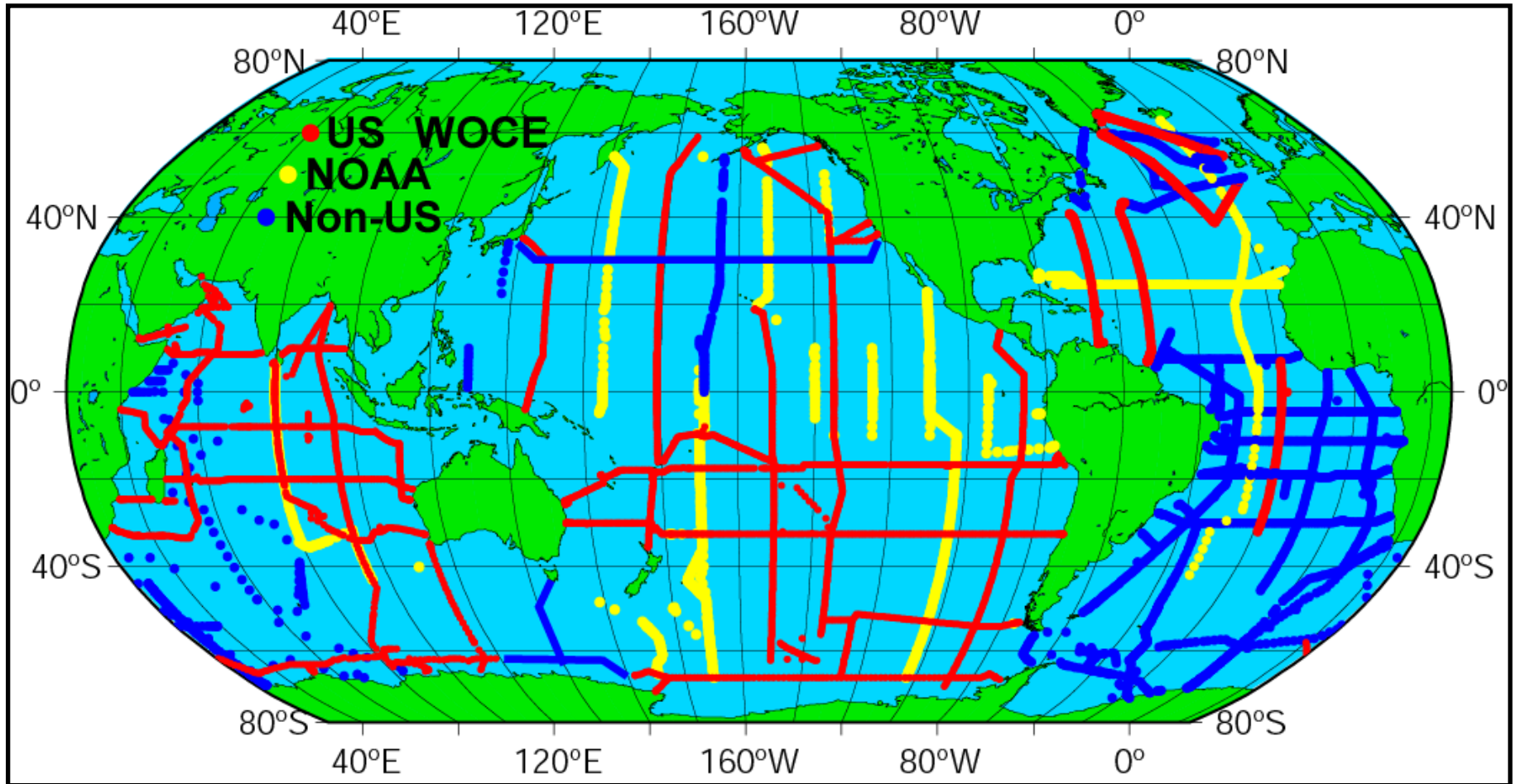


What we know about ocean CO_2 chemistry
...from time series stations



Courtesy: Richard A. Feely
NOAA/Pacific Marine Environmental Laboratory

What we know about ocean CO_2 chemistry
...from field observations



WOCE/JGOFS/OACES Global CO_2 Survey

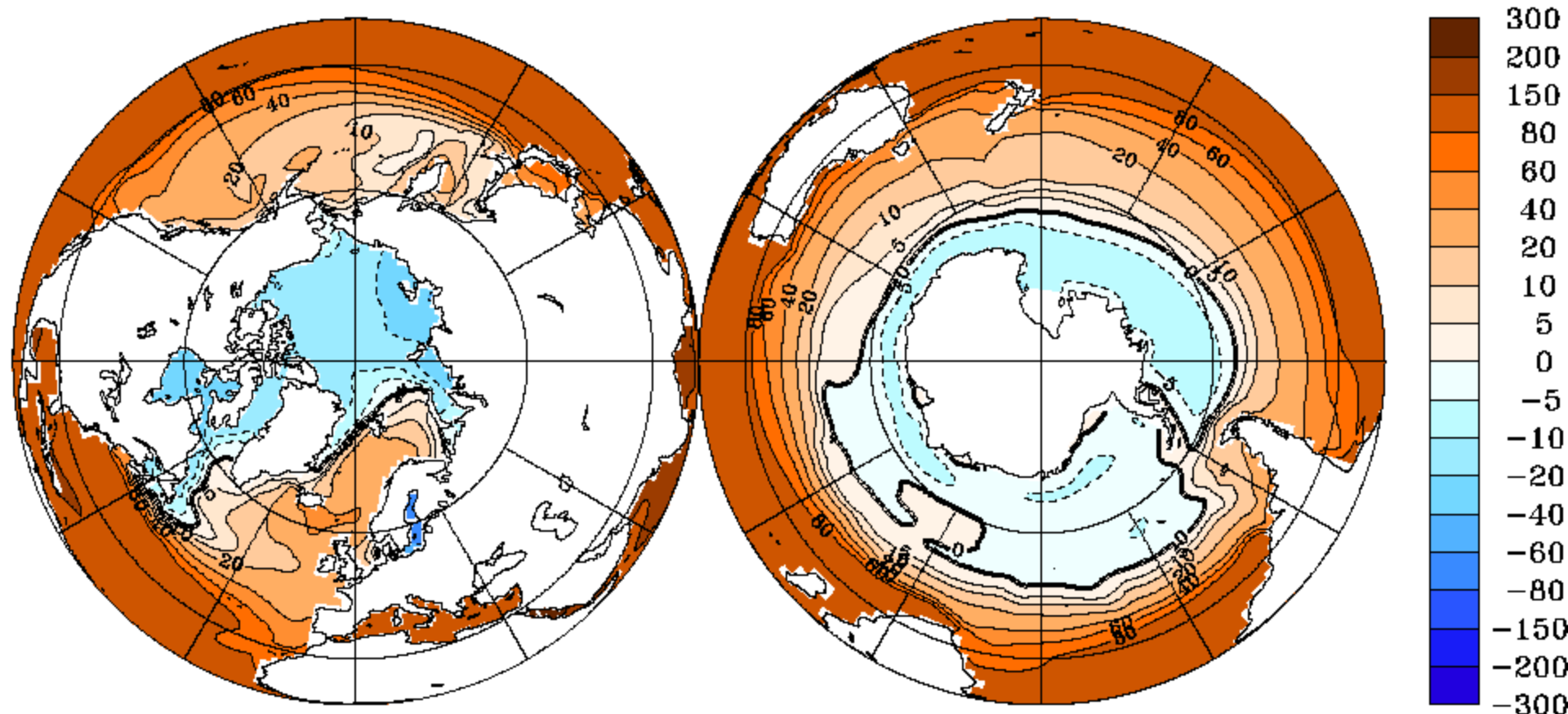
~72,000 sample locations
collected in the 1990s

DIC $2 \mu\text{mol kg}^{-1}$
TA $4 \mu\text{mol kg}^{-1}$

Sabine et al (2004)

Undersaturation is strongest in the high latitudes

Aragonite undersaturation $\Delta[\text{CO}_3^{2-}]_{\text{Arag}}$ at $2\times\text{CO}_2$



*Model approach assuming a simulation with +1% increase per year
(model results only)

Jim Orr (CEA/IAEA)

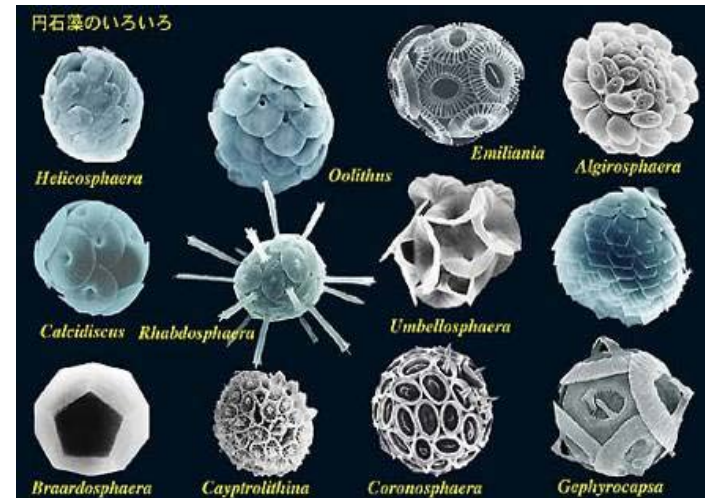
Ocean Acidification

- Decrease in pH 0.1 over the last two centuries (30% increase in acidity; decrease in carbonate ion of about 16%)
- How will this impact marine organisms and ecosystems?

Photo: Missouri Botanical Gardens



Corals



Calcareous Plankton

<http://www.biol.tsukuba.ac.jp/~inouye>

Bivalve juvenile stages can also be sensitive to carbonate chemistry



Control
 $\Omega_A = 1.5$

Hard shell clam *Mercenaria*

- Common in soft bottom habitats

Used newly settled clams

- Size 0.3 mm
- Massive dissolution within 24 h in undersaturated water; shell gone w/in 2 wks
- Dissolution is source of mortality in estuaries & coastal habitats



$\Omega_A = 0.3$

Potential impacts of high CO₂ on marine fauna

- ▶ Adverse effects on reproductive success
 - Decreased fertilization rates (sea urchins, bivalves)
 - Increased juvenile mortality (bivalves, sea urchins, copepods, fish larvae)
- ▶ Reduced growth in adults (sea urchins, bivalves)
- ▶ Impaired oxygen transport (squid)
- ▶ Reduced metabolism/scope for activity (squid)



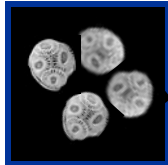
Potential Ecosystem Responses

Changes in relative abundance & distribution of calcifying species

- Non-calcifying species may outcompete calcifiers
- Geographical ranges of calcifying species may shift
- Vertical depth distributions of calcifying species may shoal with decreasing CaCO_3 saturation state

Changes in food webs and other species interactions

Potential Effects on Open Ocean Food Webs

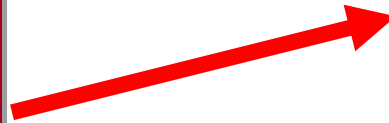
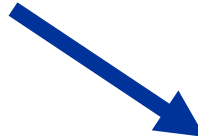


Coccolithophores



ARCOD@ims.uaf.edu

Copepods

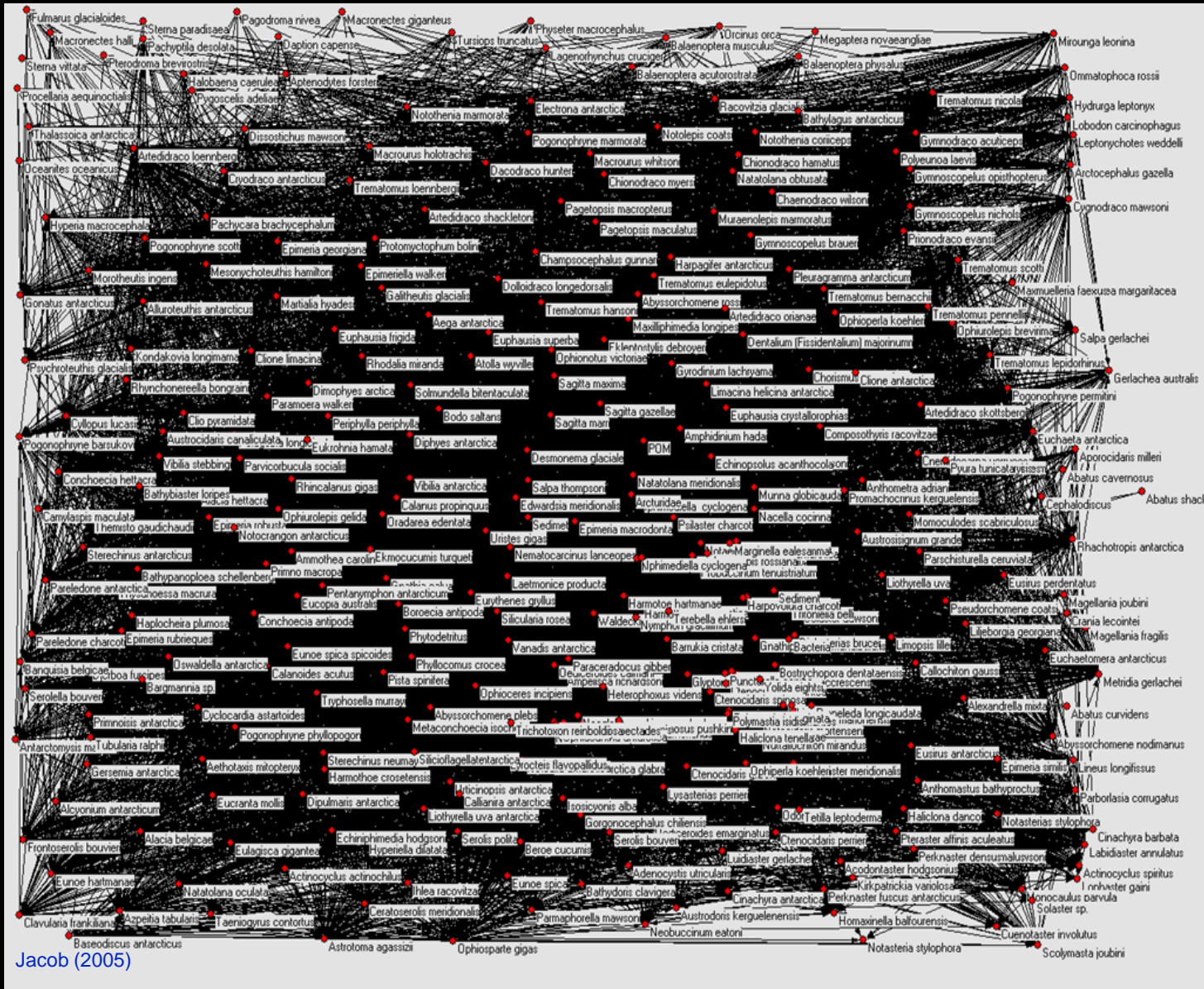


Barrie Kovish

Pacific Salmon

Vicki Fabry

Weddell Sea Food Web: 489 species (incl 62 autotrophs, >16000 trophic links (Jacob, 2005)



Potential Ecosystem Responses

Changes in relative abundance & distribution of calcifying species

- Non-calcifying species may outcompete calcifiers
- Geographical ranges of calcifying species may shift
- Vertical depth distributions of calcifying species may shoal with decreasing CaCO_3 saturation state

Changes in food webs and other species interactions

Impacts on biogeochemical cycles

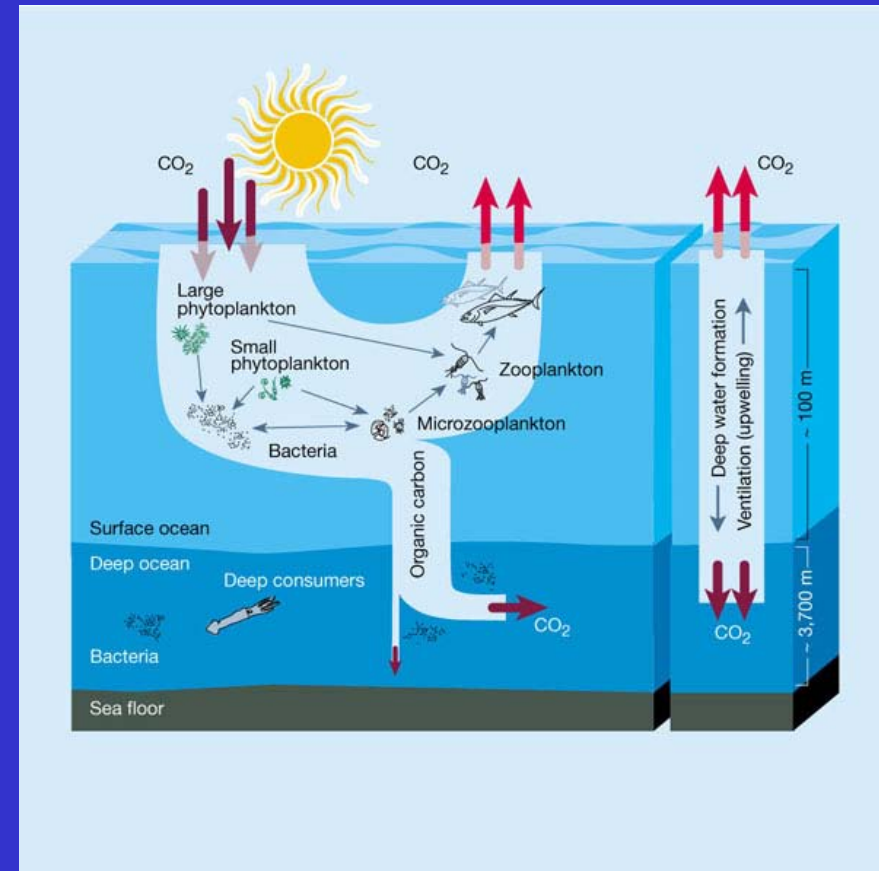
- Speciation of nutrients and trace metals
- Changes in cycling of carbon and CaCO_3 within oceans (e.g. "ballasting")
- Changes in the "microbial loop"
- Feedbacks to climate

The global carbon cycle is largely driven by biology: How will the „biological pump“ respond to OA?

What happens if biology
is turned off?

The „Strangelove ocean“:

- The biological pump stops
- The surface-deep CO_2 gradient disappears
- Within 250 yrs atmospheric CO_2 increases 2.4 times



see: Maier-Reimer, Mikolajewicz and Winguth (1996); Zeebe and Westbroek (2003)

Wrap up

- Oceans are stabilising global warming (but very slowly)
- At the same time are oceans acidifying (very fast)
- Society is facing double trouble....

Ocean acidification: a biogeological perspective

Jelle Bijma (AWI, Bremerhaven, Germany)

- Ocean acidification: present and future
- Why a biogeological perspective?
- Ocean acidification in the past
- Consequences for Biodiversity

Biological aspects

Real world

- comprises the actual complexity of the chemical, biological and ecological systems and interactions between them

Real time

- capture the time component inherent to carbonate perturbation and physiological and ecosystem response

Limitations

- gradual change makes it difficult to identify responses
- complexity of biology itself
- difficulty to capture longer term processes such as ecological adaptation, evolution and, biogeochemical cycles
- no information on recovery processes

Why paleo?

The farther backward you can look, the farther forward you are likely to see.” Winston Churchill

- What has happened *can* happen (e.g. perturbation of ocean chemistry)
- Long-term (natural) context for recent changes
- Investigate the impact on biogeochemical cycles
- Reduced complexity (integration of space and time)
- Different time scales (historical/sub-recent, G-IG, deep time,....)
- Process of recovery
- Different scenarios as case and sensitivity studies and testbeds for models

Limitations

- limited biological information (hard parts and biomarkers)
- limited by accuracy of proxy reconstructions
- restrictions on temporal and spatial resolution
- no perfect analogues

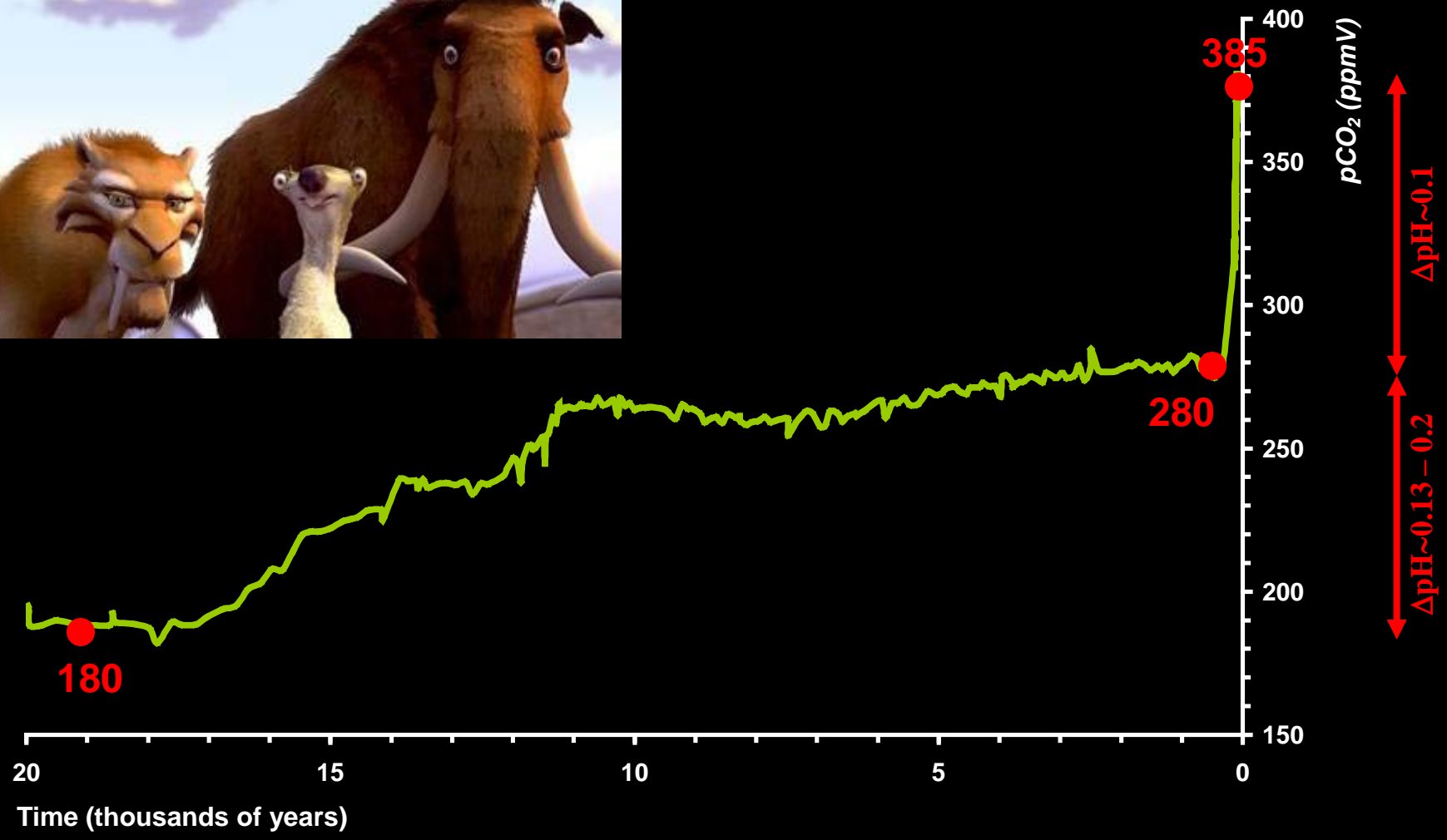
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Courtesy: Henk Brinkhuis

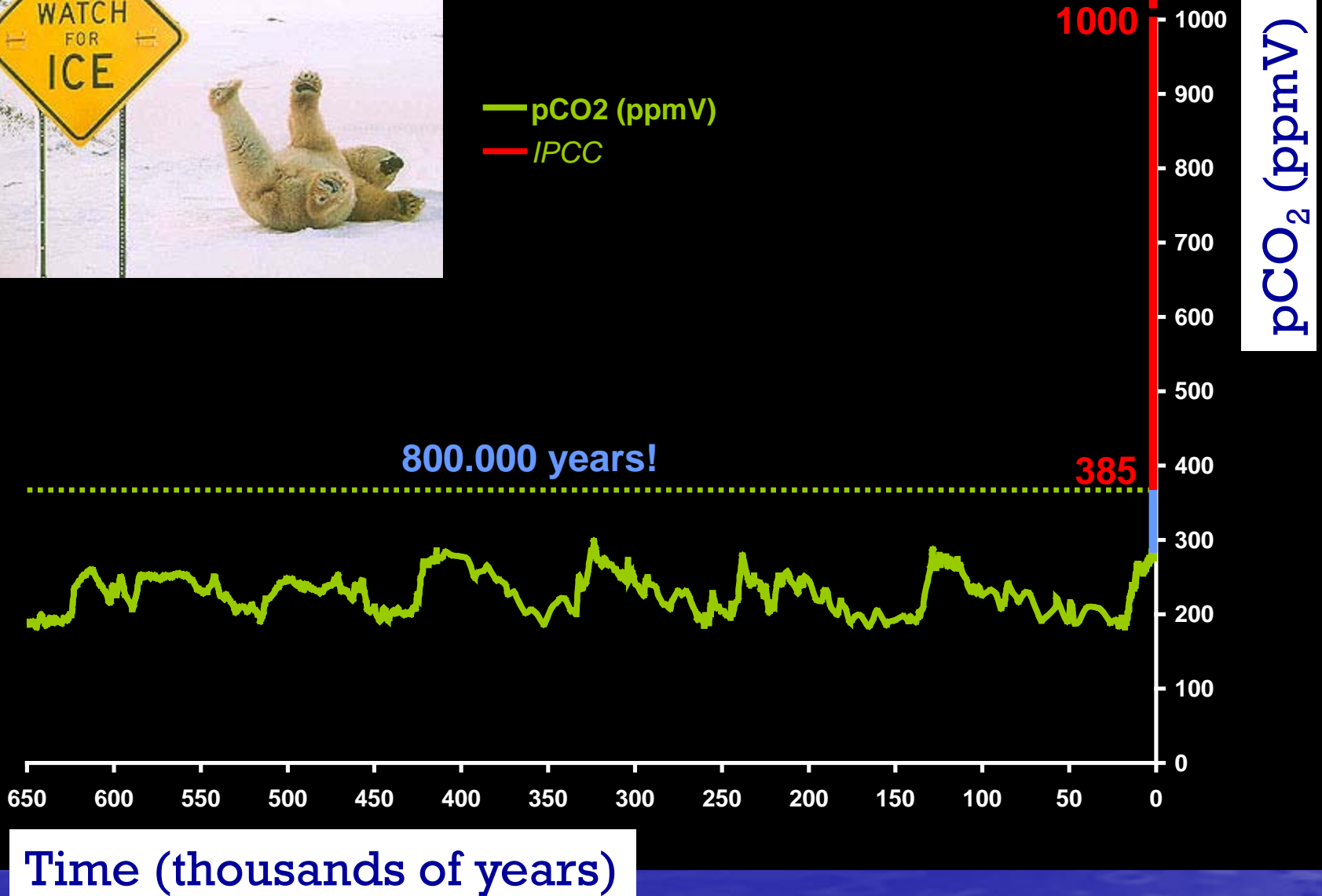
De Last Ice age



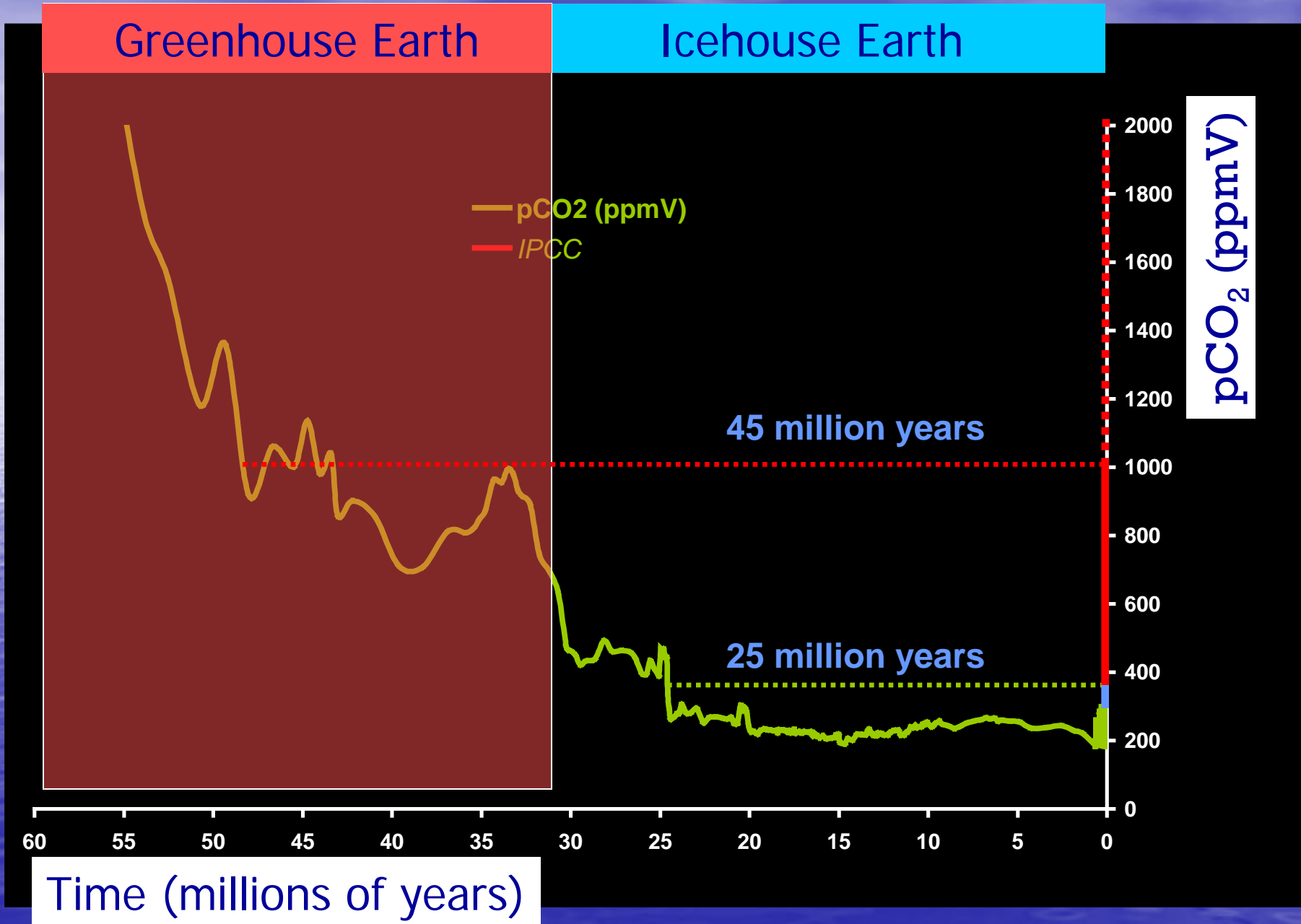
Courtesy: Henk Brinkhuis



— pCO₂ (ppmV)
— IPCC

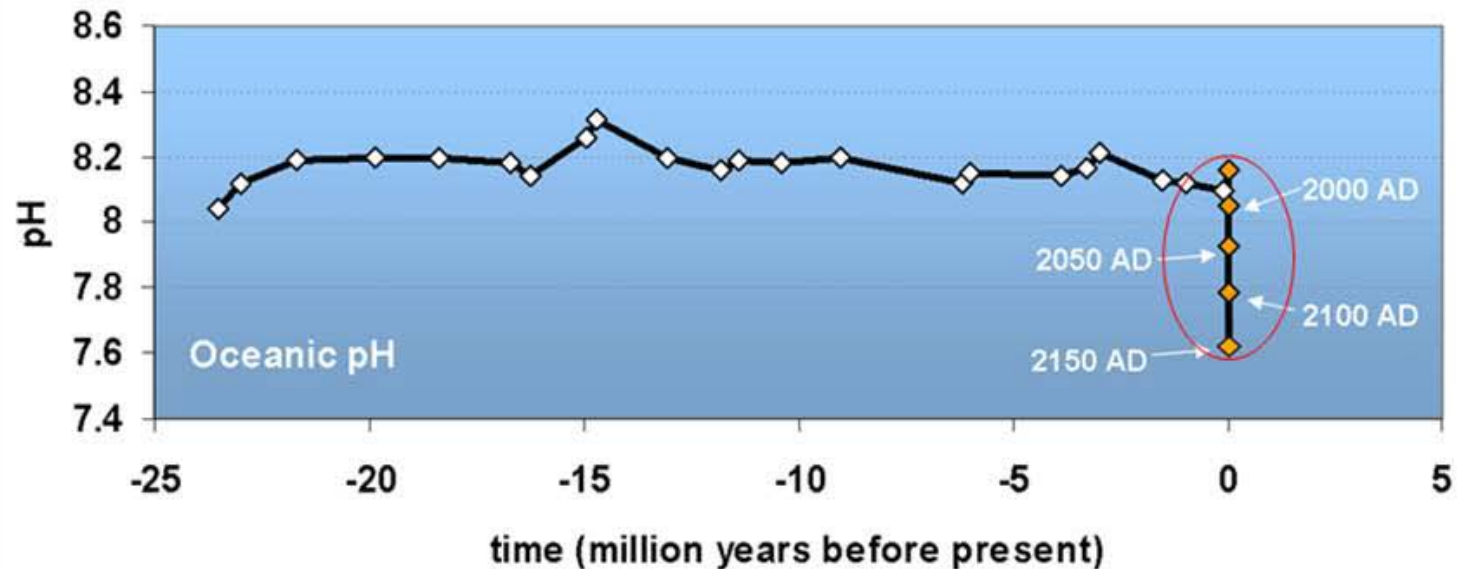


Courtesy: Henk Brinkhuis

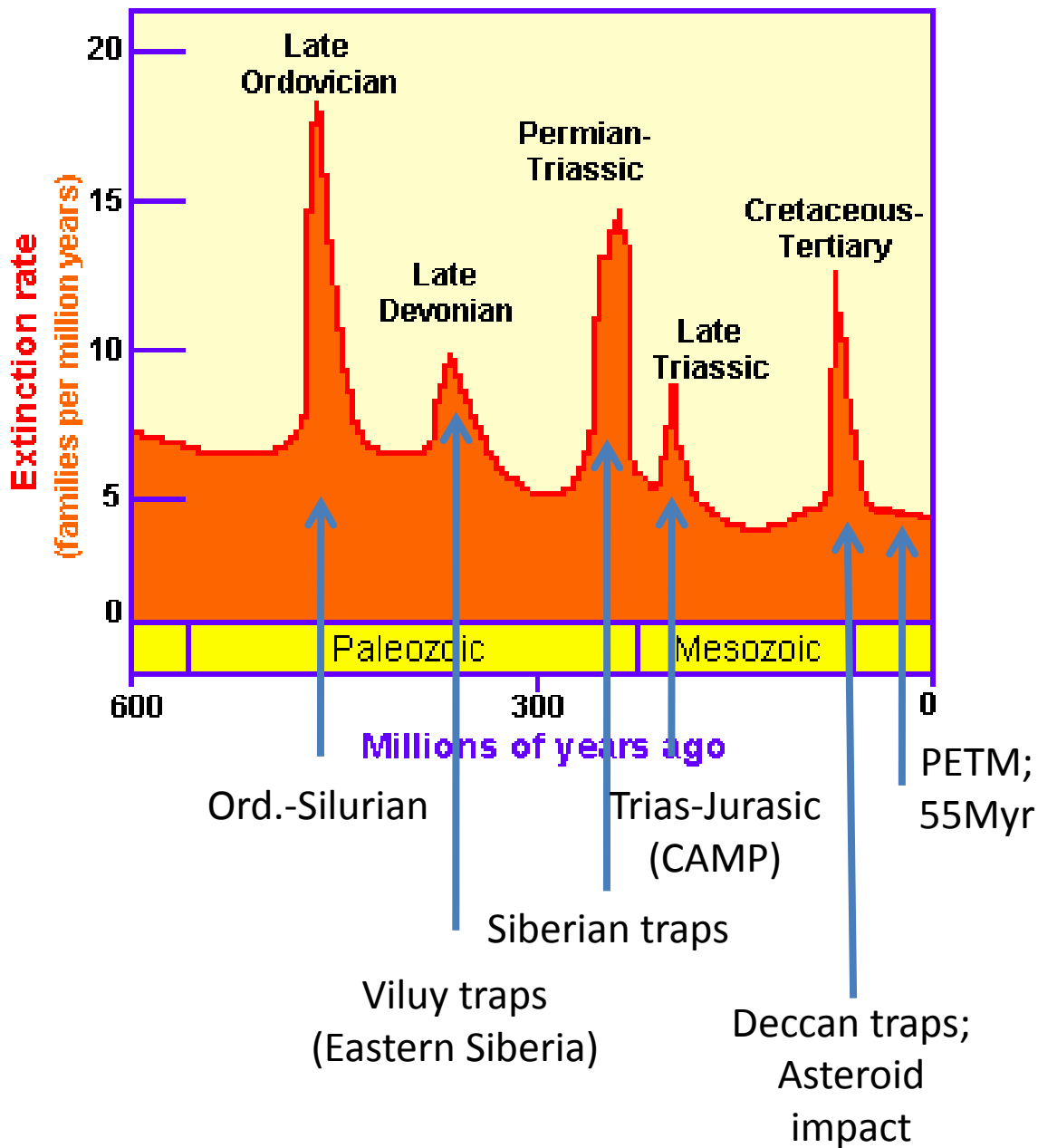


Oceans are Acidifying Fast

Changes in pH over the last 25 million years



It is happening now, at a **rate and to a level not experienced by marine organisms for ~ 20MY**



Carbon perturbation (symtoms):

- Global warming
- Ocean acidification
- Anoxia

Evidence:

- Elevated $p\text{CO}_2$ (global warming)?
- Reduced pH?
- Reduced Ω ?
- Anoxia?

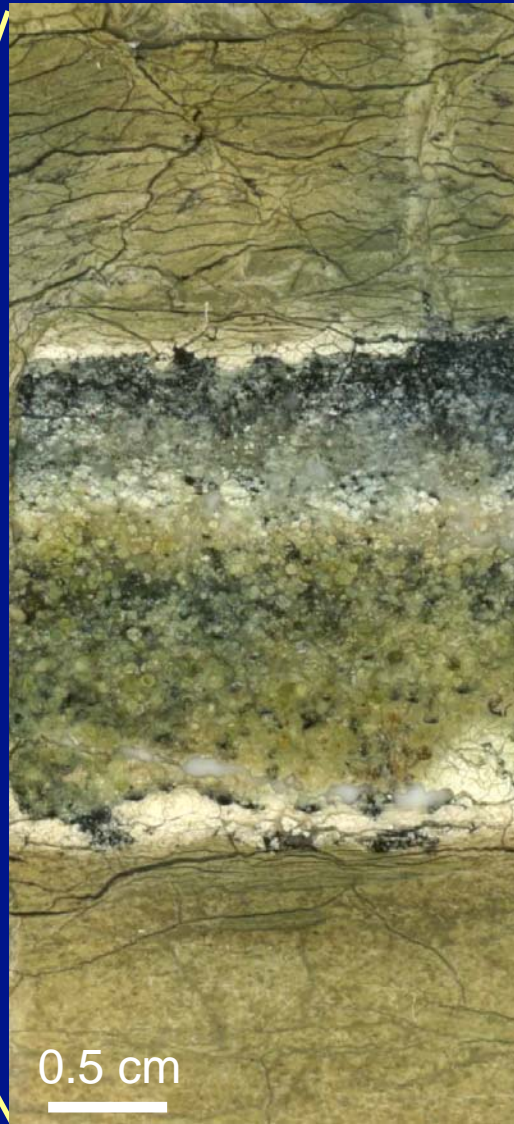


Response of marine biota to OA and climate change

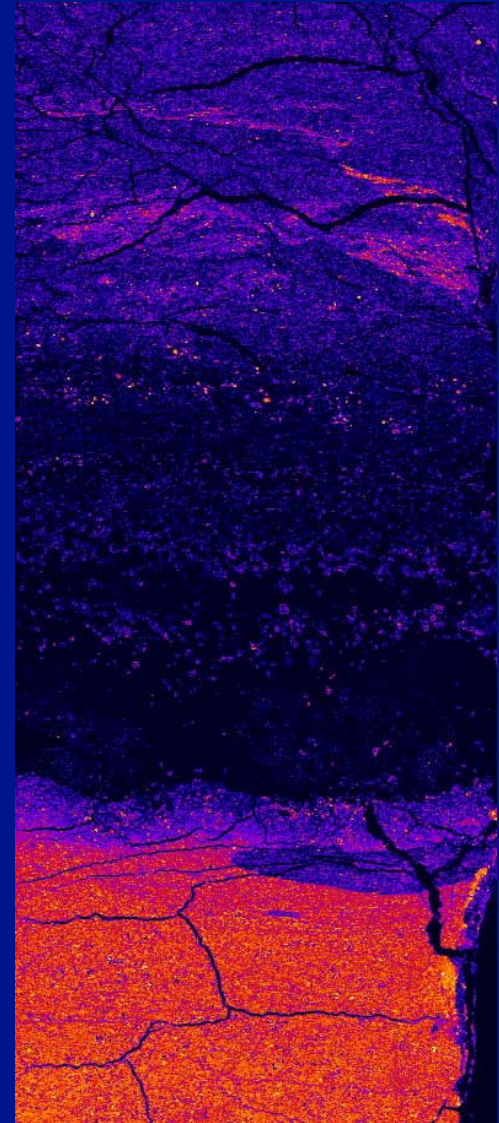
- Strong perturbation at a very fast rate → K/T impact (major planktonic extinction)

Hole 1259B 13R, 37-60 cm- boundary interval

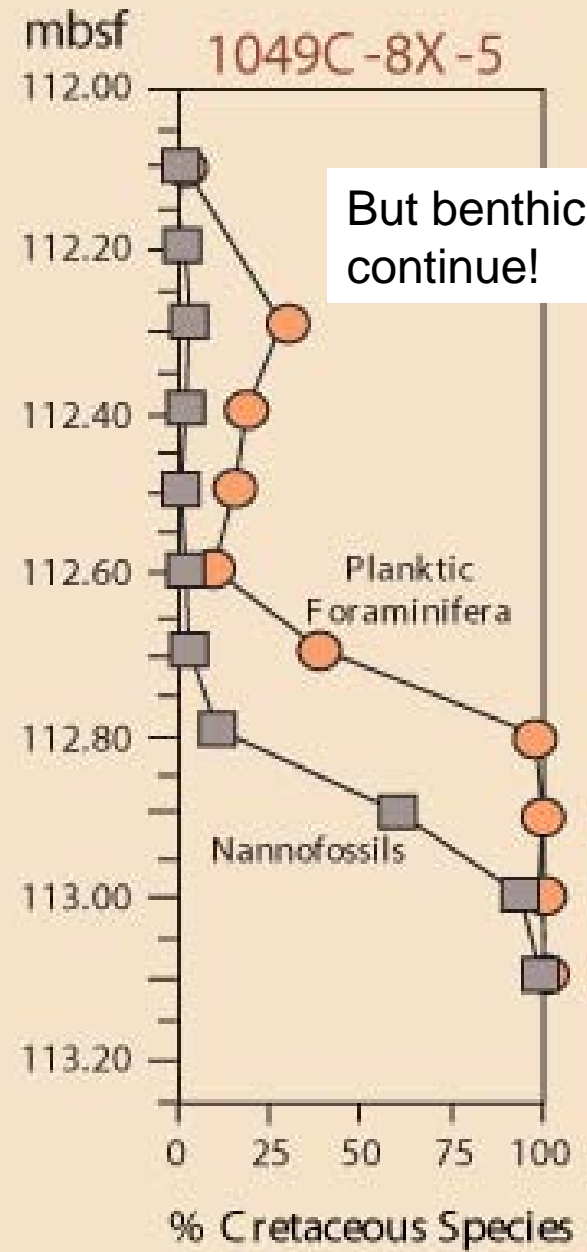
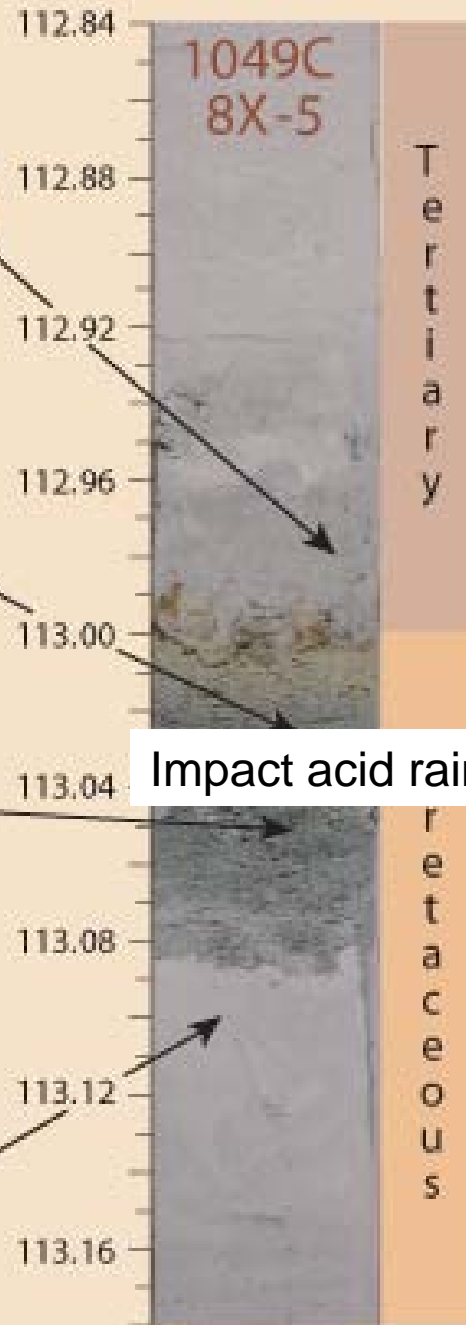
Paleogene
ejecta
Cretaceous



reflected light



Ca



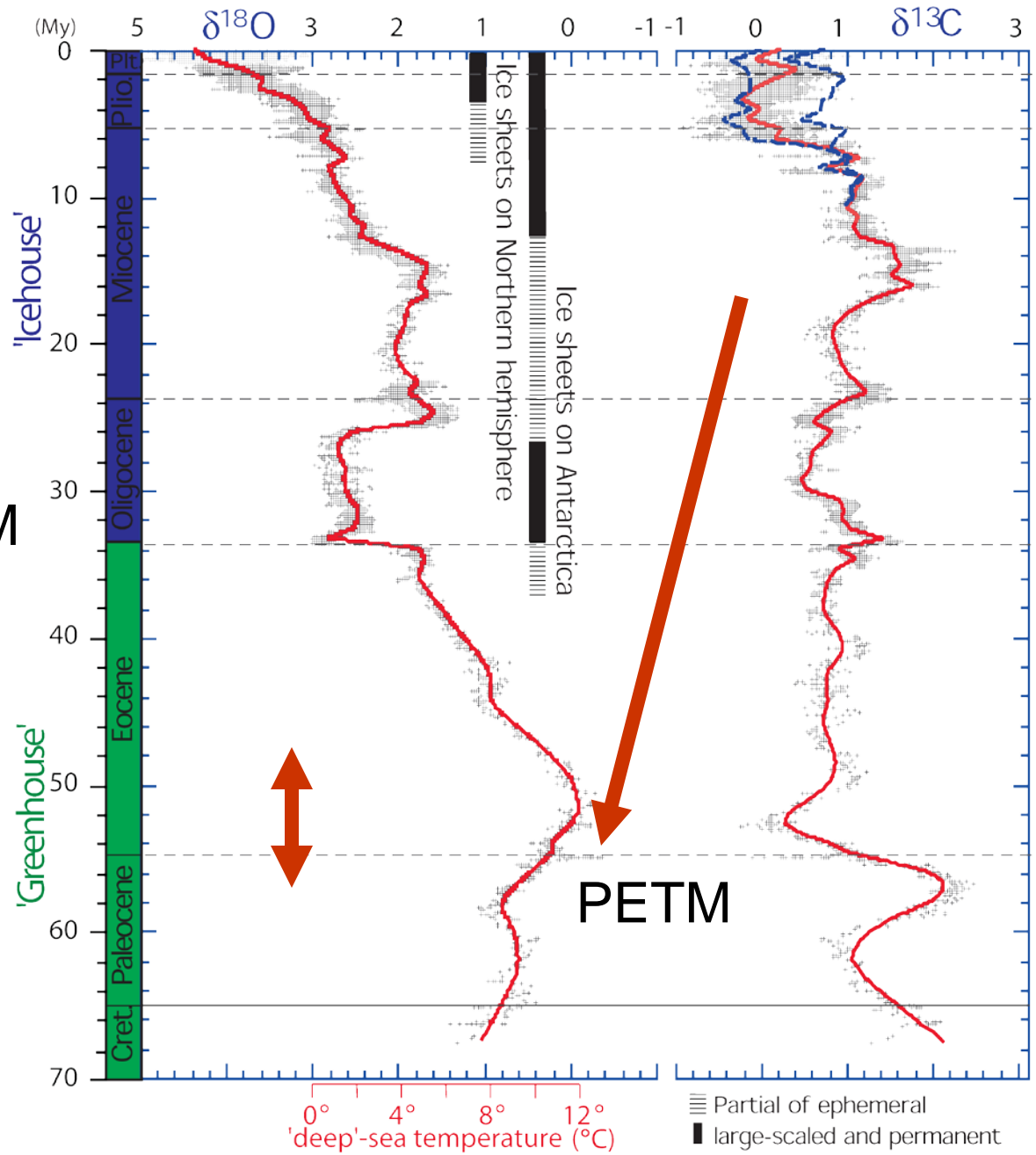
Response of marine biota to OA and climate change

- Strong perturbation at a fast rate → K/T impact (major planktonic extinction)
- Strong perturbation at a „moderate“ rate → PETM (major benthic extinction)

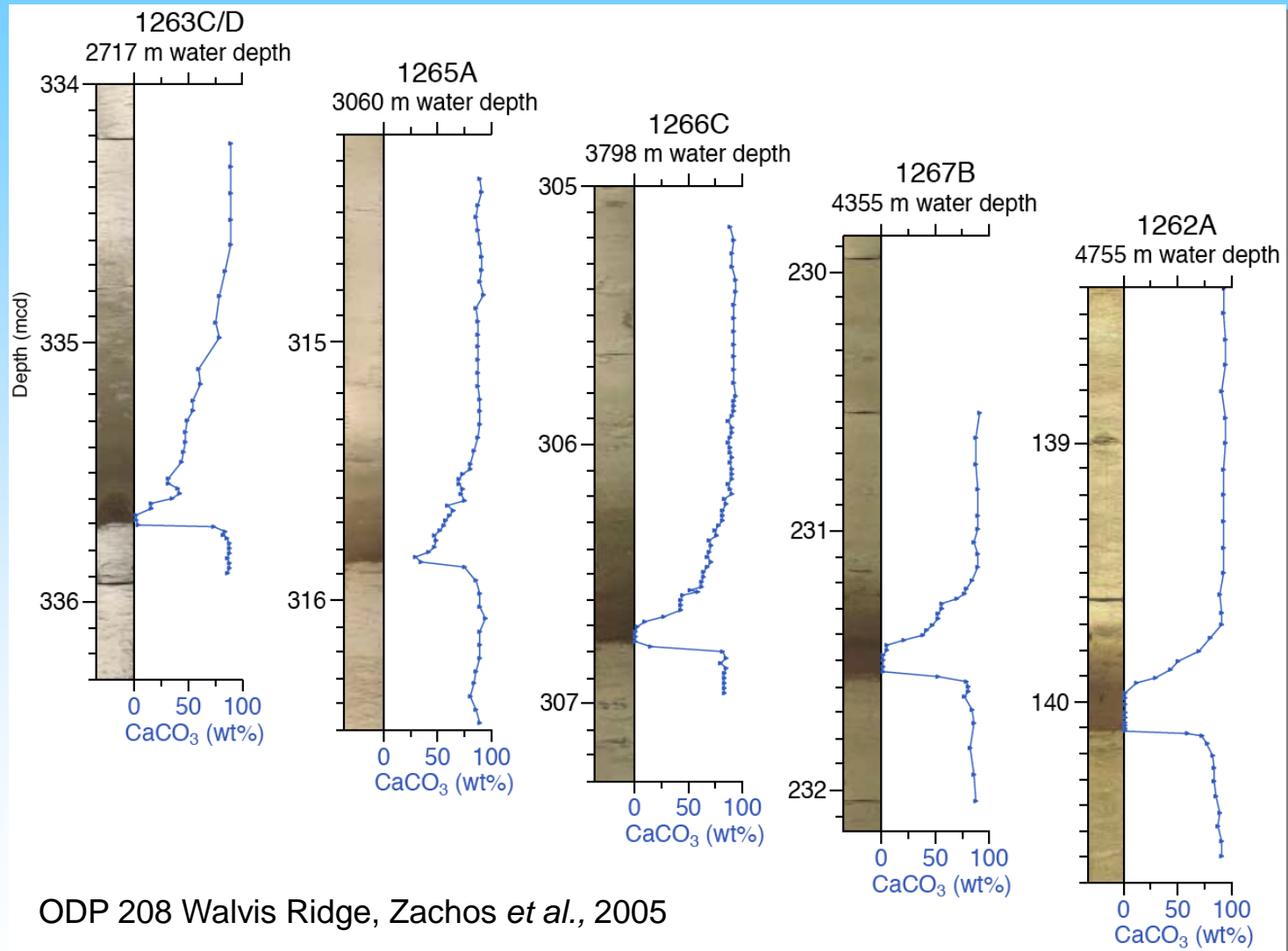
Cenozoic
long-term trends

Early Eocene: warm
superimposed: PETM

One of the largest
benthic mass
Extinction in Earth
history

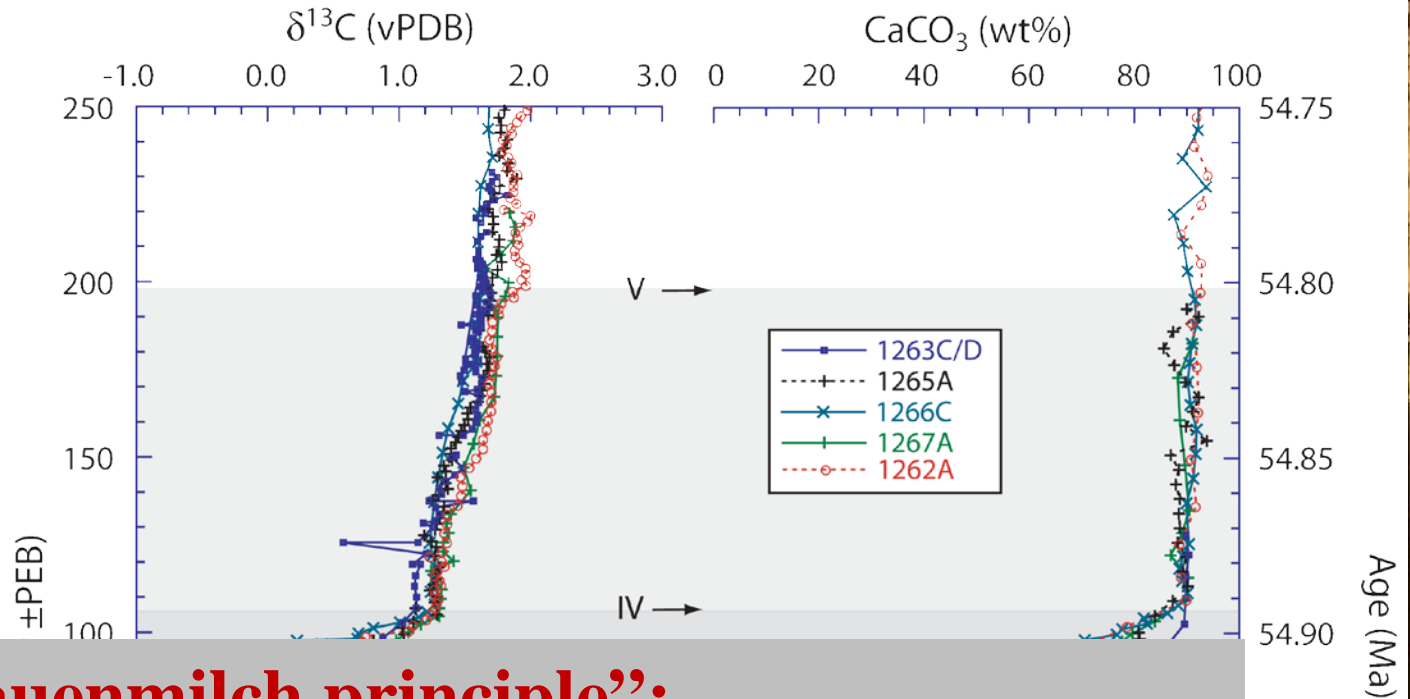


Ocean Carbonate; Walvis Ridge ODP Leg 208



„Carbonate compensation“: as lysocline is rising it destroys benthic habitats

Oceanic recovery. Walvis Ridge ODP Leg 208



“Liebfrauenmilch principle”:

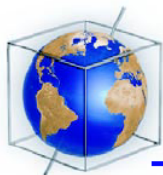
**One night of drinking followed by
2 years of hang-over**

Ca.
100ky



Response of marine biota to OA and climate change

- Strong perturbation at a very fast rate → K/T impact (major planktonic extinction)
- Strong perturbation at a „moderate“ rate → PETM (major benthic extinction)
- Small perturbation at a slow rate → Neogene, G-IG (acclimation/adaptation)



Impact of the ocean carbonate chemistry on living foraminiferal shell weight: Comment on “Carbonate ion concentration in glacial-age deep waters of the Caribbean Sea” by W. S. Broecker and E. Clark

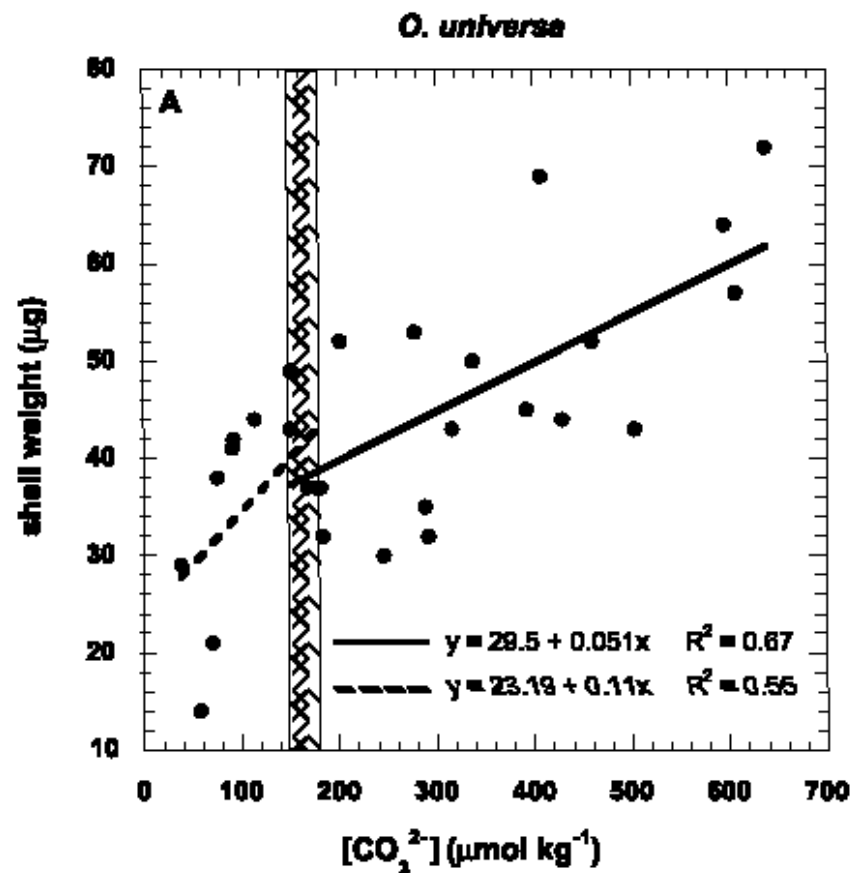
Jelle Bijma, Bärbel Hönisch, and Richard E. Zeebe

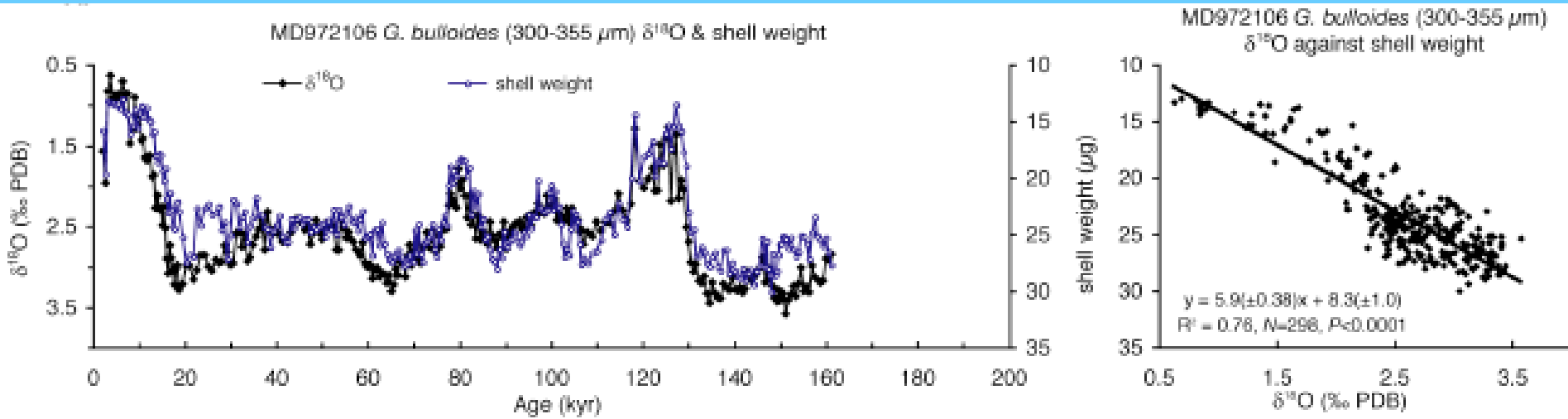
Alfred Wegener Institute, Am Handelshafen 12, Bremerhaven, D-27570, Germany

(jbijma@awi-bremerhaven.de; bhoenisch@awi-bremerhaven.de; rzeebe@awi-bremerhaven.de)

$$\Delta[\text{CO}_3^{2-}]_{\text{G-IG}} \rightarrow 100 \mu\text{mol kg}^{-1}$$

ca. 15% shell weight change



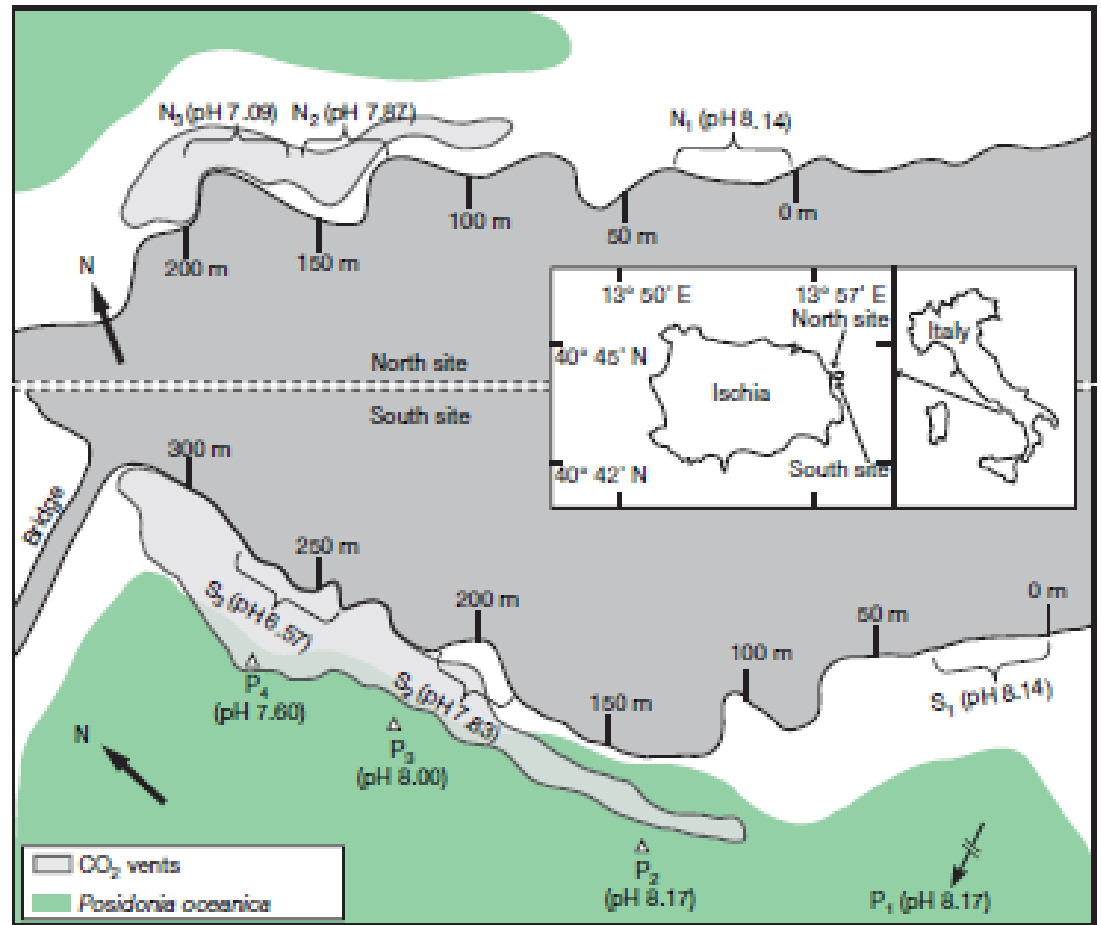


Moy, 2005

$$\Delta[\text{CO}_3^{2-}]_{\text{G-IG}} \rightarrow 100 \mu\text{mol kg}^{-1}$$

ca. 50% shell weight change!

CO₂ Vents: "Windows" into High CO₂ Ocean to Assess Ecosystem Impacts



CO₂ Vents: “Windows” into High CO₂ Ocean to Assess Ecosystem Impacts



Studies in the shallow waters of the Mediterranean and deep-sea show:

- total loss of some calcareous species
- reduced biodiversity
- “regime shifts”: totally different ecosystems

e.g. Sea grass benefit but so do invasive species



Hall-Spencer et al. Nature (2008)



Response of marine biota to OA and climate change

- Strong perturbation at a very fast rate → K/T impact (major planktonic extinction)
- Strong perturbation at a „moderate“ rate → PETM (major benthic extinction)
- Small perturbation at a slow rate → Neogene, G-IG (acclimation/adaptation)
- Strong perturb. at a fast rate → Anthropocene: decrease in species richness → breakdown of ecosystems → extinction?



.....linked to its function as a habitat and nursery for commercial fish stocks, acting as a natural barrier for coastlines, and for the provision of recreation and tourism opportunities.

A conservative value of \$100,000 km⁻² y⁻¹ (Burke and Maidens, 2004), the global economic value associated with reefs is in the order of \$30 billion yr⁻¹.

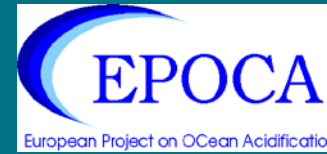


Conclusions

- ocean acidification is ongoing and future changes are very well predictable
- organismal response – poor knowledge but growing (mostly calcifiers)
- ecosystem response – not known
- evolutionary capability – completely unknown
- no perfect analogue to the present – rate of change is unprecedented
- Earth history tells us that the combination of ocean acidification, global warming and anoxia is a deadly mix

PML

Plymouth Marine
Laboratory



Marine Matters

**Is there a tipping point in
ocean acidification which
should be avoided?**



Carol Turley
Joint EPOCA, BIOACID UKOARP Workshop on Ocean
Acidification

Potential Vulnerabilities in Relation to Human Life spans

